

The bitter lakes between the past and present

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The Bitter Lakes of Suez Canal were characterised with layers of salt deposits of about 13m thick. After the opening of the Suez Canal in November 1869, a salt barrier of salinity greater than 65 ‰ was formed in the lakes which acted as an obstacle for the migration of the marine organisms between the Mediterranean and the Red Seas.

Four cruises were conducted in the Bitter Lakes during summer periods (1982-1986), to study water circulation, decrease of salinity with time, and situation of the salt barrier in the lakes. Temperature, salinity and currents were measured at different depths. The mean water level during the same period was also obtained.

A minimum salinity of about 43.5 ‰ has been recorded during the last few years. A difference in salinity values of about 1.0 ‰ is observed between winter and summer seasons in comparison with that obtained in 1966 (4.0 ‰). The water in the Great Bitter Lake circulates in clock- and anti-clock wise directions conformable with the tidal currents. Due to the prevailing north winds in summer, the water piles up in the southern part of the Great Bitter Lake, which generates an anticyclonic motion in the vertical plane as well as drives a southward current depending on the amount of water flowing from Lake Timsah, Lake Menzalah and the Mediterranean Sea.

After the last deepening and widening of Suez Canal and Bitter Lakes in 1976, the evaporation potential is highly significant factor in increasing the salinity of the lakes during summer. Such increase is estimated as 1.2 ‰. Finally, the salt bed is about to be exhausted and its effect on salinity is insignificant. The salt barrier which has been dominating for more than hundred and twenty years is going to be disappeared. Presently, the migration of the marine organisms between the joined seas can occur occasionally at any time without any osmotic problem.

Dynamical modal analysis for the Coastal Seas around Turkey

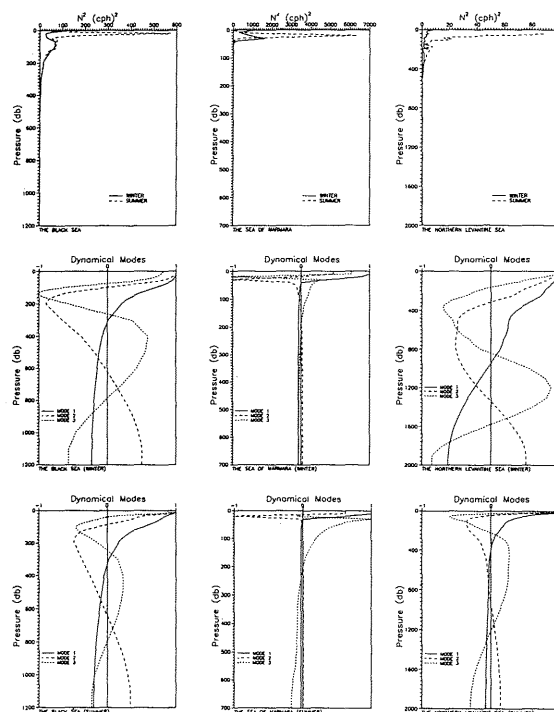
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The hydrographic data collected in six surveys of R/V Bilim in the Black Sea (BS), the Sea of Marmara (MS) and the Northern Levantine Sea (NLS) during 1988-1989 are used to perform the vertical modal analysis. Utilizing the Brunt-Vaisala frequency (N^2) and geopotential anomaly profiles obtained from averaged density field over the region, the relative contribution of the dynamical modes to the observed field of motion are determined by solving the vertical structure equation for each cruise.

The computed N^2 profiles show that the BS, MS and NLS are characterized by distinctly different stratification properties. In the BS, both the winter (April 1989) and summer (September 1988) N^2 profiles display presence of a two-layer stratification comprising the upper layer of 150m with substantial density variations and the layer beneath possessing almost no variations below 300m. The summer and winter N^2 profiles only differ within the uppermost 20-30m in response to seasonal thermohaline forcings. The stratification observed in MS for the winter (March 1989) and summer (September 1989) reveals even stronger two layer structure with a sharp pycnocline at 25m separating two water masses. We note that the values of N^2 is one order of magnitude greater than those for the BS. In the NLS of the Mediterranean basin, the winter (March 1989) and summer (July 1988) N^2 profiles have substantial differences. The summer case reveals a surface layer of about 40m and subsequent transitional layer extending down to 80m with significant N^2 variations. The winter case, on the other hand, possesses no seasonal variations and is governed by relatively small density differences in the water column as implied by small values of N^2 .

Modal analyses of the data show that the amplitudes of the first baroclinic mode in the Black Sea for both summer and winter seasons are confined to the surface layer and have zero crossings at 300m. The first baroclinic Rossby radius of deformation (R_1) is found to be 19 km in each season. The first and third baroclinic modes respectively contain 80% and 10% of the total available potential energy (TAPE) in summer although the first mode contains 92% of the TAPE in winter. In NLS, R_1 is determined as 8 km and 11 km for the winter and summer cases, respectively.



The first baroclinic mode has a zero crossing at 400m and the second at 50m and 900m for the summer case. The first mode contains 90% of the TAPE. The winter analysis, however, reveals substantially different modal structure. The zero crossing of the first mode occurs at 950m and the second at 320m and 1320m. The first mode contains only 82% of the TAPE. 14% of it is found in the second baroclinic mode. The MS exhibits completely different modal structure. R_1 is determined as 15km and 20 km for the winter and summer cases, respectively. The TAPE is partitioned as 50% and 48% between the first and second baroclinic modes in the summer case. This partition occurs between the first three successive modes as 32%, 38% and 20% for the winter case.