Relative Sea-Level Oscillations and Depositional Patterns on the Ebro Distal Continental Margin : Plio-Quaternary Evolution

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The Plio-Quaternary deposits of the Ebro distal margin are developed over a major boundary, Reflector M of Messinian age. This reflector was eroded when most of the margin was subaereally exposed due to the sea level drops of the Messinian salinity crisis. The sedimentary sequences, above Reflector M, define four major environments (shelf, slope, base-of-slope, basin floor) corresponding to several systems tracts related to major sea level oscillations. Lowstand and highstand systems tracts are differentiated on the basis of reflectors characteristics, unit geometry and interpreted depositional environment. Information from piston cores samples and boreholes from the oil industry and DSDP Site 122 complement the identification of depositional units, and allow for a time framework to correlate with global eustatic sea level oscillations (Ryan, Hsü et al., 1973; Garcia-Sifieriz et al., 1979; Alonso and Maldonado, 1990; Dañobeitia et al., 1990).

1990; Damobeitia et al., 1990). Lowstand systems tracts are characterized basinward by shelf margin deltas near the shelf edge, thick slope muds on the slope, and channel-levee complexes, base-of-slope aprons and interchannel deposits at the base-of-slope region. These deposits, accounting for most of the stratigraphic thickness of the Quaternary sequence show a stratigraphic continuity of selsmic reflectors between these two provinces that may represent a contemporary deposition (Fig. 1). The most active growth periods of the distal margin occurred during intervals characterized by relative sea-level fails, similarly to siliciolastic turbidite deposits throughout the geological record. Slope deposits are largely related with shelf-edge spillower and distal prodeita high density flows on the upper slope, with gravity-driven nepheloid flows on the middle slope, and with turbidite and mass-flows on the lower slope (Baraza, 1989). The base-of-slope deposits were also developed during these periods of low-sea stands and they are associated with two slyles of deposition: (1) flushing of sediments from river discharge, outer shelf and upper slope environments trough slope canyons, to base of slope channel-levee complexes, and (2) with unchannelized mass flow processes, from slope mass-failure in areas of unstable slope terrain resulting in the base-of-slope aprons (Alonso and Maldonado, 1990). The highstand systems tracts are characterized at the slope and base of slope by benipelagic processes over most the distal margin, although some transparent, high energy facies of the base-of-slope aprons may also occurs. These deposits were developed during the latest rising eggment of the eustatic curve, high sea level stand and initial sea lowering (Fig. 1).

Begment of the technicity four epochs in the growth patterns of the distal Ebro margin during the Plio-Quaternary (Fig. 1): (A) during the "salinity crisis" the emergence of the Ebro margin allows the carving of extensive erosional surfaces and the by-passing of sediments to the distal basin plain; (B) the Pliocene flooding of the Mediterra-nean Sea after the Messinian time, and generalized global eustatic high sea level favoured the development of depositional units characterized by shallow marine facies in the fluvial valleys and a drape of fine-grained deposits (Ebro Clays) over most of the continental margin, corresponding to a <u>hightand systems tract</u> (C) the generalized lowering of sea-level of the Late-Pliocene displaced the sediment supply from Ebro sequence, and initiates a <u>lowestand systems tract</u>; and (D) the subsequent sustatic-ci matic oscillations of the Late Pliocene-Quaternary time slope wedges, channel-levee complexes and base of slope aprons, alternating with lowe



slope wedges, channel-levee complexes and base of slope aprons, alternating with low energy, sediment drape units that represent respectively low and highstand systems tracts. * <u>low</u> tracts

Figure 1. Lowstand (LST) and highstand (HST) systems tracts and four epochs (A, B, C, D) in the growth patterns of Ebro distal continental margin.

REFERENCES

ALÓNSO, B. AND MALDONADO, A., 1990. Late Quaternary sedimentation patterns of the Ebro turbidite systems (Northwestern Mediterranean): two styles of deep-sea deposition. In: C.H. Nelson and A. Maldonado (Eds), Northwestern Mediterranean Sea. Mar. Geol., (in press).

BARAZA, J., 1989. Procesos de edificación y características geotecnias del talud continental del Ebro. Doctural Thesis, U.P.C., 471 pp.

DANOBEITIA, J.J., ALONSO, B. AND MALDONADO, A., 1990. Geologica framework of the Ebro continental margin and surrounding area. In C.H. Welson and A. Maldonado (Eds), Northwestern Mediterranean Sea Mar. Geol., (in press). 1990. Geological In:

GARCIA-SINERIZ, B., QUEROL, R., CASTILLO, F. AND FERNADEZ, J.R., 1979. A new hydrocarbon province in the western Mediterranean. 10th World Petr. Congr., Bucarest, p. 191-197.

RYAN, W.B.F., Hsu, K.J. et al., 1973. Initial reprts of the De Drilling Project, XIII. U.S. Print. Office, Washington. 1447 p ep Sea

Rapp. Comm. int. Mer Médit., 32, 1 (1990).