A Preliminary Study of the Principal Recent Sediment Types along the Eastern Margin of the Aegean Sea

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Fourty-six surficial sediment samples collected during the oceanographical cruise of the R/V Billm in Eastern Aegean Sea in 1987, from depths ranging between 12 and 640 m (Fig. 1), were subjected to granulometric, carbonate, organic carbon, and optical investigations. The main goal is to provide data that contribute to increase the knowledge on sedimentary processes in this poorly known part of the Aegean Sea. Preliminary results suggest that - on the basis of biogenic CaCO₃ content - four main types of Recent sediments overlay the floor of the Eastern Aegean Sea, namely very low-calcareous (< 5 % CaCO₃), low-calcareous (5-25 % CaCO₃), calcareous (25-50 % CaCO₃), and high-calcareous (50-75 % CaCO₃) sediments. Carbonate components of sediments are largely made up of benthonic-shelly materials mostly occurring in the surroundings of GOKgread Island-Strait of Dardanelles, south of Bozcaada Island, Chios Island-Çeşme Peninsula, and Bodrum Peninsula-Kos Island.



Most of the organic carbon values of sediments (0.30-0.70 %) are comparable with those found elsewhere in the Eastern Mediterranean (ERGIN et al., 1988), but somewhat lower than those from the Sea of Marmara (ERGIN and EVANS, 1988). The organic carbon contents of the Aegean Sea sediments of this study, in general, reflect low biogenic production here, compared to sediments from highly productive Marmara waters. Exceptionally high organic carbon concentrations (up to 3.50 %) occurred at or near river mouths, and also at sites with high sea-grass communities. Among the biogenic components, the occurrence of calcareous algae "Lithothamnium" is characteristic, especially off the coast of Bozcada Island and in areas of between Kos Island and Bodrum Peninsula. Sediments are composed of materials ranging in grain size: from silty clay to gravelly-muddy sand. Mud is widely distributed off the river mouths (areas of high terrigenous input) and in the embayments (areas of low energy conditions), particularly in Edremit and Izmir Bays, as well as, the Karaburun Peninsula-Lesbos Island-Chios Island triangle.

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Geophysical Framework of the Sea of Marmara

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Western Turkey has undergone dominantly N-S extension causing the formation of E-W trending grabens and similarly oriented normal faults (Eyidoğan, 1988). Marmara Sea is the extension of the Trakya basin in the north. Necogene lateritic depositions are present at the borders of this basins. The islands which are also called Marmara, are made up of crystalline Paleozoic rocks of marbles and granites. The Caspian-Black Sea region up to Italy in the west is at present tectonically very labile. Rapid subsidence characterizes a series of more-or-less elongated basins run subparallel to and are interrupted by areas undergoing uplit. Thus, during the Quaternary, the Sea of Marmara has subsided an amount well in excess 1000 m, accompanied by extension and transform motion. On the other hand, Quaternary uplit in wester Anatolia is widespread. The early Quaternary roesional plateau has raised to up to 300 m in Trakya, 950 m on Tekirdağ and almost 2000 m on Uludağ in the south. These uplifts and subsidences are the extensive phenomena traceable to regional tectonics. During the final phase of the destruction of the Tethys, readjustments of the microcontinents and collision lead to the development and reactiviation of transform, as well as non-rigid deformation of continental crust along the collision front. As a result, north-south crustal shortening and secondary east-west stretching occurred. The primary vertical tectonics is uplift, but where transform faults cut across the direction of transform motion, crustal attenuation and subsidence prevail, leading to the formation of pull-apart basins. transform motion, cr of pull-apart basins.

Marmara region has different seismic characteristics from the rest of western Marmara region has different seismic characteristics from the rest of western Anatolia and appears to act as a separate tectomic unit (Crampin and Erans, 1986). This region shows higher seismic activity than the western Turkey in general indicating that this region is partly under the influence of the western end of the North Anatolian Fault which splays into a number of branches in and around the Sea of Marmara (Dewey and Sengór, 1979). The northernmost branch becomes a graben and follow the Gulf of Izmit, connecting the Cinarcik pull-apart basin in the Sea of Marmara (Seneth et al. 1985). Bread en full perchaption existing a contrained for the set of the senether and the set of the senether and the set of the set o (Sengor et al., 1985). Based on fault mechanism solutions, Crampin and Evans (1986) suggested that the Marmara block is beging rotated and sheared in order to accommodate the right-lateral motion of the North Anatolian Faault and extensional tectonics of the southwestern Anatolian province.

The structures of the Sea of Marmara show the characteristics of rapid The structures of the Sea of Marmara show the characteristics of rapid subsidence acconpanied by extension and transform motion. In this area, pure strike-slip motion changes into extensional strike-slip movement responsible for the creation of the basins of the Sea of Marmara and the North Agean where the Ganos Dağ area active faults joins these two basins. It might be well possible that the Saros trough and the Sea of Marmara basin (Adatepe, 1988; Ergûn et al., 1988) were activated after the extension ceased in the north where the Trakya basin is situated having 3 to 4 km of Neogene sediments.

The seismological, seismic reflection, gravity and magnetic geopysical data will be reviewed for the Sea of Marmara and the surrounding area. Earthquake data indicate seismolgically very active region. The estimated crustal thickness is around 30-35 km determined from the earthquake data. The northern side of the Sea of Marmara shows smoother gravity and magnetic anomalies than the southern side. Magnetic anomalies are very much affected by the magmatic end volcanic intrusions on the southern side of the Sea of Marmara. The basement (most probably made up of the crystalline rocks df the Biga Peninsula on the south and the Istranca in the north) has been obtained to be around 2 to 5 km determined by the gravity modelling and power spectrum analysis. Neotectonic movements show their effects even within the very recent sediments indicated by shallow seismics in the area. Also, there exists a crushed zone in the middle of the basin marking the effects of strik-silp motion coupled with tensional movement. Izmit Bay area in the east of the Sea of Marmara was found to be a hall graben from the gravity interpretation with the vertical movements taking place on the south side. movements taking place on the south side.

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