## Climatic variability in the Northeastern Mediterranean and Black Sea

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The general circulation of the Eastern Mediterranean has been the object of many research efforts. It is known to be dominated by important basin and sub-basin features such as gyres, jets, eddies, meandering currents reflecting its complex geometry, bathymetry and highly variable atmospheric forcing. Its seasonal and higher frequency variability and their regional dependence have been studied in sufficient details in number of studies (see Hopkins, 1978; Malanotte-Rizzoli and Hecht, 1988; Ozsoy et al., 1989). Yet, almost nothing is known about the eventual existence and importance of interannual variability, on time scales of decades or so, of the general circulation of E. Med. The recent interest on climatic oscillations and trends brings this subject to actuality, since the oceans should respond to changes in atmospheric forcing.

In this paper, we examine the low frequency variability, on time scales of decades or so, of coastal sea surface temperatures (SST) in the northeastern Mediterranean and Black Sea, and their eventual relation to atmospheric temperature (AT). Changes in SSTs, especially if they are not totally coherent all over the area, can be considered as an indication of changes in the density field, and therefore circulation.

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The long-term variabilities were studied using monthly means of STT and ATS from 20 trations covering the Adriatic, Tonian, Aegan and ATS from 20 trations covering the Adriatic, Tonian, Aegan spikes and unrealistic values. The common data set was formed for the period 1955 - 1984. The largest variance in both SST and AT time-series is contained in the annual cycle. In order to study longer term variability, the annual variations and other high-frequency signals having periods shorter than two years were eliminated by means of the 24m214 filter (Thompson, 1983).

The resulting time-series do show the existence of quite important variations of SST and AT inside the area during the study period. The main features of these variations are:

a.-Both SST and AT time-series in the Aegan and southern Adriatic Sea show a constant trend of decrease (0.5-1.0 deg.C.) from the series of the constant of the series of th

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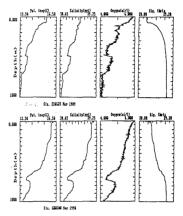
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## Large scale inversions in the North Levantine Basin

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Measurements carried out by R/V Billm during 1988-1990 in the north east Levantine basin show inversions associated with large scale intrusions of anamolus water masses at depths between 500-900 m. The thickness of the intrusions is between 100 m and 200 m, and they are observed primarily in the fronts between the Rhodes cyclonic gyre and the adjacent anticyclonic eddies. The measurements were taken with a SBE-9 CTD continuous profiling system. The temperature and salinity values of the intrusions differentiate them from the well known Levantine intermediate water (LIW). Both warm and cold intrusions have been observed. The warm water intrusions have relatively higher oxygen than the ambient water; the opposite is true for the cold intrusions. Examples of a warm and a cold intrusion are shown below. At station E30.25 (N34°30, E29°25), the intrusion layer is located between 500 m and 700 m; its temperature and salinity are higher by 0.2°C and 0.08 ppt. than the ambient. The intrusion at station GOOL30(N36°, E29°30) extends between 600 m and 750 m, and its temperature and salinity are lower by 0.3°C and 0.1 ppt than the ambient. The potential density profile shows that the intrusions are hydrostatically stable.



The temperature, salinity and the oxygen content indicate that the cold intrusions result from lateral sliding of the upwellwing water of the Rhodes gyre towards the periphery of the gyre. The sources of the warm intrusions are the two, anticyclonic gyres centered approximately at  $N34\,^\circ\!\!\!/\, E30^\circ\!\!\!/$ , and at  $N36\,^\circ\!\!\!/\, E30^\circ\!\!\!/\, 20$ 0. While the main intrusion layer is stable, Brunt Vaisla frequency plots show that the interface regions are unstable, implying a gradual mixing between the two water masses.