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The seismic stratigraphy of the central region in the Gulf of Cadiz exhibits three majo sequences well characterized in both single and multichannel reflection profiles (Fig. 1). Th underlying sequence forms the nucleus of linear diapir structures intersecting the slope in WSW-ENE direction. The intermediate sequence is composed of several units wit discontinuous and parallel reflectors, separated by unconformities. These units are largell attributed to low-energy, terrigenous deposits, which are characterized by the absence of major channels or erosional/depositional features. The youngest sequence is Pilo-Quaternary in agand it shows at the base a major erosional unconformity, which is deeply entrenched in the underlying deposits. This sequence is made of contourite facies which formed on the easter Gulf of Cadiz continental slope. Here the Mediterranean undercurrent has flower northwestward parallel to the slope contours since the Pliocene opening of the Strait of Gibraltar. Gibraltar

northwestward parallel to the slope contours since the Pliocene opening of the Strait of Gibraltar.

The contourite facies change westward and downstream because of an interaction betwee the linear diapiric ridges that are perpendicular to the slope contours and the progressiv northwest decrease in speeds of the undercurrent. Coincident with the decrease i undercurrent speeds from the Strait of Gibraltar, the following northwestward gradation of contourite seismic facies occurs: (1) sand dune contourite facies on the upstream mid-slope terrace, (2) sediment-drift facies banked against the diapiric ridges, and (3) smooth slope facie with generally continuous, parallel-stratified reflectors (Fig. 2). Downstream, there are severa hundred-meter thick sediment drift wedges of sediment drape deposits and high-ener deposits with irregular erosional surfaces and broad-scale, low-angle unconformities.

In the upstream contourite facies, sonar images reveal a wide variation of bedform field on the present seafloor, caused by variations in undercurrent speeds, as well as by changes is superficial sediment texture. These bedforms range laterally from (a) 2-D, transverse sand dunes in channel floors, to (b) 3-D barchan dunes that laterally evolve to (c) large, straight crested 2-D dunes, and to (d) 3-D irregularly-shaped dunes, in the silt-covered slope trace. It contrast, down-valley Mediterranean undercurrent ribbons flowing along the channel between the ridges modify the basic east to west sequence of contourite facies. The down valley facies exhibit major erosional truncation, and extensive cut and fill facies on seismi profiles (Fig. 1). Sonographs collected across the channelized central sector show (e) erosiona carps, sand patches or rock-outcrops exposed on channel floors; (f) small and parallel 2-I dunes, or (g) regularly shaped, variably size, 3-D dunes along valley walls.

The surface contourite and sediment drift facies on the Gulf of Cadiz slope have formed undercurrent Late Pleistocene hemipelagic-drape facies

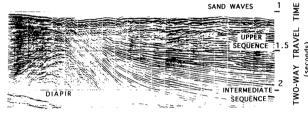


Fig. 1. - High resolution MCS profile from sand wave area showing the three main hoseismic sequences in the central sector of the Gulf of Cadiz.

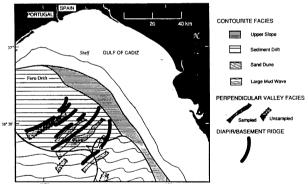


Fig. 2.- Distribution of contourite and other current-deposited facies of the Cadiz margin slope (Modified from NELSON et al., 1992).

REFERENCES

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