## Tectonic Facies Concept Applied to Interprete the Geology of Cyprus

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Mountains formed by plate collision are exemplifed by the West Alps. Rock formations of the colliding plates are (1) overthrust, (2) underthrust, or (3) pushed out onto foreland. The three modes of deformation correspond respectively to those of (1) the Austro-alpine rigid basement nappes, (2) the Penninic melanges and mobilized basement nappes, and (3) the Helvetic cover nappes of the Swiss Alps. HSU and BRIEGEL (1991) proposed to consider the Austro-Alpine, Penninic, and Helvetic three major tectonic facies, and interprete the genesis of less well-known mountain belts of the world through comparison with the Alps. with the Alps.

The present Benioff zone under the Hellenic Arc dips to the north, and the Neogene tectonics of the eastern Mediterranean is commonly considered to have a southward vergence. Extrapolating from the Holocene to Miocene, the structural deformation of Cyprus has also been assumed to be a southward continuation of structures in Anatolia, characterized by a southward vergence; the Troodos ophiolites has thus been considered a fragment of ocean crust thrust up- and southward along one or more north-dipping thrust faults (e.g., BIJU-DUVAL and MONTADERT, 1978; ROBERTSON, 1990; 1991).

Two major tectonic units are present on the Island of Cyprus. The Kyrenia Range is underlain mainly by upper Mesozoic and Cenozoic formations, forming disharmonic folds and decollement thrusts. This style of deformation is typically that of a peeled-off passive-margin sequence. The southern half of the island is underlain mainly by the Troodos margin sequence. The southern hal Ophiolite and by the Monia Melange.

That the Troodos represents ancient ocean crust is a concensus interpretation; the ophiolite complex is considered a huge slab in a tectonic melange. That the Monia Melange represents a tectonic melange now sandwiched in the suture zone between colliding plates is also a concensus interpretation. The question is the vergence of the deformation. Assuming that the Benioff Zone south of Cyprus dipped to the north during the Miocene as the Benioff Zone under the Hellenic Arc, the orthodox interpretation is assume that a micro-continent, called by ROBERTSON (1990; 1991) the Monia Microcontinent has been underthrust beneath Cyprus. The orthodox model assumes thus that the Kerenia formations belonged to an overriding block.

The theoretical model for collision tectonics predicts, however, that a tectonic melange is the accretional wedge on an active margin, formed by shearing within a Benioff Zone which dips under the overriding block; the Saas-Zermatt and Arosa ophiolite melanges dip under the Austro-Alpine nappes. The Kerenia style of deformation is, however, not comparable to the rigid-basement thrusting of the Austro-Alpine nappes; the style is more analogous to the tectonics of the Helvetic nappes. Considerations of comparative tectonics require that the Kerenia is a foreland deformed belt, formed by the plunger action of a *traineau écraseur* of an overriding plate.

HSU suggested in a memorandum to JOIDES drilling in 1989 that the Miocene Benioff Zone south of Cyprus may have dipped south, not to north. According to this model, the Troodos ophiolite underlies a back-arc basin which had its origin in Late Cretaceous back-arc seafloor-spreading. The Upper Cretaceous-Tertiary sedimentary sequence of the Kerenia Range is that of a passive-margin behind, i.e., north of the backarc basin. The frontal arc should then be located under the Eratosthenes Seamount! Such a paleogeographic interpretation suggests that the tectonics of the Cyprus deformation is typically that of a back-arc basin collapse. The ocean lithosphere under a back-arc basin, like that under the South China Sea today, dips under a frontal arc, causing the elimination of the basin and an arc-continent collision. The suture zone of such a collision is located on the back side of the arc, exemplifed by the Lichi Suture of the Mio-Pliocene arc-continent collision in Taiwan. The Monia Melange was the accretional wedge under the sandwiched in the suture. The plunger action of the overriding arc and the melange caused the decollement deformation of the Kerenia-Range sequence.

This interpretation of the Cyprus tectonics presents an alternative to the current orthodoxy. Ocean- drilling south of Cyprus should yield data to discriminate the two working hypothesis. If the vergence of the Miocene deformation was directed southward, drilling south of Cyprus should penetrate a foreland deformed belt, similar to that of the Helvetic Alps, or that of the Appalachian Valley and Ridge. If, however, the vergence of the deformation caused by the Miocene collision was directed northward, ocean drilling should penetrated an island arc sequence (volcanic or non-volcanic) under the upper Neoene sediments of the Eratosthenes Seamount.

The settlements of the fundational setundation. The settlement of the question of the Cyprus tectonics might help resolve the controversy on the tectonics of the Antalya Melange. RICOU and others (1979), the "super-nappists" suggested a southward thrust of the ophiolites from northern Turkey during the suturing of Anatolia. ROBERTSON (1990; 1991), on the other hand, considered the Antalya Ophiolite the lithosphere under a small ocean basin, the Isparta Angle; the ophiolite, according to this model, was emplaced "outward" from the Isparta Angle and onto the adjacent, relatively autochthonous carbonate platforms." My interpretation of the Cyprus geology supports the postulate by ROBERTSON. In the scheme of a back-arc basin complex, the continental crust under the Kerenia Range (and under the Troodos Ophiolite) could be interpreted as the foundation of a remmant arc, similar to the KyushuPalau Rideg of West Philippine Sea. The Kerenia passive-margin sequence was laid down on the south side of this remnant arc. That Antalya and Troodox ophiolites were, according to this model, the ocean lithosphere under the two back-arc basins north and south of the remnant arc respectively; those basins were eliminated by the mechanism of back-arc basin collapse, and the ophiolite nappes were thrust northward onto carbonate platforms.

## REFERENCES

BIJU-DUVAL B. & MONTADERT L., 1978.- Introduction. In BIJU-DUVAL B. & MONTADERT L. (eds), The structural history of the Mediterranean Basins, Paris: Editions Technip, p. 1-12.
HSU K.J. & BRIEGEL U., 1991.- Geologie der Schweiz. Basel: Birkehauser, 219 pp.
RICOU L.E., MARCOUX J. & POISSON A., 1979.- L'allochthonie des Bey Daglari orientaux. Bull. Soc. France, 21: 125-133.
ROBERTSON A.H.F., 1990.- Tectonic evolution of Cyprus. In J. Malpas et al. (eds.) Ophiolites, oceanic crustal analogues, Proc. Symposium Troodos, 1987. Geol. Survey, Dept. Cyprus, pp. 235-252.
ROBERTSON A.H.F., CLIFT P.D., DEGNAN P.J. & JONES G., 1991.- Paleogeographic and paleotectonic evolution of the Eastern Mediterranean Neotethys. Palaeogeography, Palaeoclimatology, Palaeoecology, 87: 289-343.

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