

Canyons, chenaux profonds et appareils détritiques provençaux et corses, après les campagnes Mesea I et Mesea II (Mesim)

Gilbert BELLAICHE\*, Guy PAUTOT\*\*, Laurence DROZ\*, Virginie GAULLIER\*

et équipe scientifique embarquée

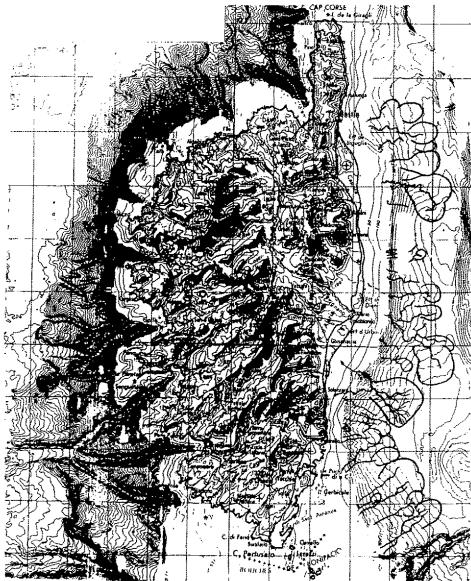
\*Lab. de Géodynamique sous-marine INSU-CNRS, Observ. Océanol.

VILLEFRANCHE-SUR-MER, (France)

\*\*Département des Géosciences marines. Ifremer, Centre de Brest, PLOUZANE, (France)

En l'espace de deux années, deux campagnes océanographiques entreprises dans le cadre du programme national de reconnaissance des zones économiques de l'Ifremer, ont permis de compléter entièrement la couverture bathymétrique détaillée au sondeur multifaisceaux et sismique des marges continentales provençale et corse. La première s'est déroulée en 1990 à bord du Jean Charcot et au moyen du Seabeam, la seconde, en 1991, à bord de l'Atalante et au moyen du Simrad EM12.

Les résultats acquis font l'objet d'un travail de cartographie systématique au centre Ifremer de Brest (cf. G.PAUTOT et G.BELLAICHE, ibid). Ils permettent en premier lieu de compléter et de préciser de façon notable nos connaissances sur la morphostructure de ces marges et notamment sur les appareils sédimentaires de bas de pente. Ainsi, la précision des sondages, permet de pouvoir rattacher désormais sans ambiguïté ces appareils à leurs canyons vecteurs respectifs et à leurs différentes sources d'alimentation terrigène. Ces résultats fournissent également des données nouvelles, notamment sur les marges corses, dont les caractères dissymétriques sont nettement mis en évidence. Alors que la marge occidentale de cette île présente des caractères très voisins de sa partie conjugée provençale, sa marge orientale se particularise par la présence, dès le rebord de la plateforme, d'un très grand nombre de petits appareils détritiques particulièrement actifs, intensément ravinés, isolés ou coalescents, en relation parfois directe avec le vigoureux réseau hydrographique continental. Au niveau des bouches de Bonifacio, les mouvements tectoniques qui ont présidé à l'ouverture de la mer Tyrrhénienne commencent à se faire sentir, comme en témoigne le cours segmenté des canyons de ce secteur méridional.



1- Outre les auteurs cités, l'équipe scientifique embarquée au cours de ces deux campagnes comprenait notamment J. R. Vanney, A. Coutelle, J.C. Aloisi, C. de Giovanni, J.P. Rehault, J. Deverchere, J.P. Maze et S. Monti.

Saline minerals of the Tuzla salt deposit as indicators of Paleoceanographic conditions

V. BERMANEC<sup>1</sup>, D. TIBLJAS<sup>1</sup>, M. CRNJAKOVIC<sup>2</sup> and G. KNIEWALD<sup>3</sup>

(1) Depart. of Mineralogy and Petrology, Fac. of Science, Univ. of Zagreb, ZAGREB (Croatia)

(2) Croatian Museum of Natural History, Mineralogical Section, ZAGREB (Croatia)

(3) Center for Marine Research Zagreb, "Rudjer Boskovic" Institute, ZAGREB (Croatia)

The Tuzla rock salt deposit in the central part of Bosnia and Herzegovina (Yugoslavia) comprises the principal salt body of the Tusanj hill area, and the recently discovered Tetima salt-stock lens. The salt-dome type deposit is largely stratified and is hosted in a Miocene-age sedimentary formation consisting primarily of banded marls with anhydrite. The salt bearing sedimentary sequence belongs to the Majevica mountain range horst, prominent feature of a Miocene sea archipelago. Inspite of earlier detailed investigations of the salt deposit, there exist as yet no unequivocal evidence regarding the geological origin of this formation. This is mostly due to the fact that event lake basins such as the Tuzla area have an inherently low preservation potential for the geological record (HUTCHINSON, 1957). However, recent studies of trace element distributions in an assemblage of lithotype indicator minerals imply a possible complex marine type or mixing-zone depositional model of the evaporite formation (KNIEWALD *et al.* 1986; BRAJKOVIC *et al.* 1988).

The term "saline mineral" as used in this study refers to all minerals associated with the principal evaporite series. Most of these satisfy the condition of having sodium in one or all of their cation sites, with the carbonate anionic group supplying all or part of the negative charge. Other accessory minerals, such as the iron sulphide group minerals, are also treated here due to their close association with some minerals of the "saline" suite.

**Principal minerals**

The principal mineral paragenesis of the evaporite series consists of halite, thenardite and anhydrite, the  $\alpha(\text{H}_2\text{O})$  indicator pair (cf. BRAITSCH, 1971) being thenardite-mirabilite. The other classic indicator pair gaylussite-pirssonite is missing entirely from this assemblage.

**Accessory minerals**

Apart from the three principal minerals, the evaporite series assemblage contains also a number of accessory (supporting) minerals present occasionally in trace quantities. These are bradleyite, glauberite, huntite, mirabilite, nahcolite, northupite and ferringorthupite, probertite, searlesite, tseelite and trona. Iron sulphides are closely associated with northupite, and are in some cases enclosed within crystals or globules of this mineral. The northupite-ferringorthupite couple is highly sensitive to prevailing conditions of redox potential (EH), either during the time of initial deposition or in the course of diagenesis. Dolomitic limestones are found closely related to evaporite beds. Evidence of progressive dolomitization involves the calcite-huntite-dolomite equilibria and may account for their formation under evaporative, non-evaporative or seepage reflux.

The depositional milieu of the saline mineral suite can be analysed and interpreted in terms of the following matrix of environmental parameters (adapted from KELTS, 1988):

HYDROLOGY	FACIES	CONCENTRATION	IONIC COMPOSITION
Open	Shallow	Dilute	acidic
Semi-closed	Littoral	Oligosaline	$\text{Ca}, \text{CO}_3$
Closed	Deep	Mesosaline	$\text{Na}, \text{Mg}$
	Pelagic	Saline	$\text{Cl}, \text{SO}_4$
		Penosaline	$\text{Na}, \text{HCO}_3, \text{CO}_3$
		Hypersaline	$\text{Mg}, \text{Ca}, \text{Cl}$

Even a preliminary analysis of the associate mineral assemblage indicates that the evaporite sequence and the hosting sedimentary series was probably formed in an environment of rapidly changing physico-chemical conditions ranging from a coastal marine setting to a shallow, hypersaline lacustrine brine. There is also evidence for periods of suboxic/anoxic and euxinic regimes, which corroborates the mixing-zone model for the formation of the Tuzla salt deposit.

**REFERENCES**

- BRAITSCH O., 1971. - Salt deposits. Their origin and composition. Springer Verlag, Berlin, 299 p.
- BRAJKOVIC Z., BERMANEC V., TIBLJAS D. and KNIEWALD G., 1988. - On the origin and geological type of the Tuzla salt deposit in Yugoslavia. 2. Trace element geochemistry of lithotype indicator minerals. *Rapp. Comm. int. Mer Medit.*, 31/2, 95.
- HUTCHINSON G.E., 1957. - A treatise on limnology. John Wiley & Sons, New York, 342 p.
- KELTS K., 1988. - Environments of deposition of lacustrine petroleum source rocks. *Geol. Soc. Special Publication No. 40*, 3-26.
- KNIEWALD G., BERMANEC V. and TIBLJAS D., 1986. - On the origin and type of the Tuzla salt deposit in Yugoslavia. 1. A trace element study of northupite and halite. *Rapp. Comm. int. Mer Medit.*, 30/2, 72.