

OCEANIC DATA ASSIMILATION IN THE MEDITERRANEAN SEA

P.G. DRAKOPOULOS¹, M. COOPER¹, K. HAINES¹, A. LASCARATOS², P. WU¹

¹ Department of Meteorology, University of Edinburgh,
King's Buildings Edinburgh, EH9 3JZ U.K.

² Department of Applied Physics, University of Athens, 33 Ippocratus Str.,
Athens, 10680 Greece

At University of Edinburgh a new technique has been developed for projecting altimetric data to produce deep ocean currents. The method is based on physical and dynamical conservation laws which ensure that water mass properties such as temperature, salinity and potential vorticity are preserved on isopycnal surfaces which remain below the mixed layer. These methods have been successfully tested on a simplified version of an eddy resolving Cox model for the Atlantic Ocean (HAINES, 1994). In this work we present results of applying this technique in a more realistic Cox model for the Mediterranean, using a twin experiment approach. The model which is eddy producing, has $0.25^\circ \times 0.25^\circ$ horizontal resolution and 19 vertical layers and is forced seasonally (ROUSSENOV *et al.*, 1994). The assimilation run started 20 years after the spinup began and was integrated forward in time for one year. During assimilation time (every 10 days), the model density profiles were displaced vertically in response to observed surface pressure anomalies, in such way to allow geostrophic currents to decay with depth, thus avoiding unrealistic barotropic changes. Errors introduced in the density, temperature, salinity and velocity fields, mainly due to mesoscale eddy activity in the Levantine basin, the Alboran Sea and along the north coast of Africa have been successfully reduced at the end of the run (e.g. Fig. 1). Salinity and temperature errors were reduced by 50% and velocity by 40%. In addition, and despite the ventilation of certain isopycnals during water formation periods, potential vorticity error on these layers was also found to be improved.

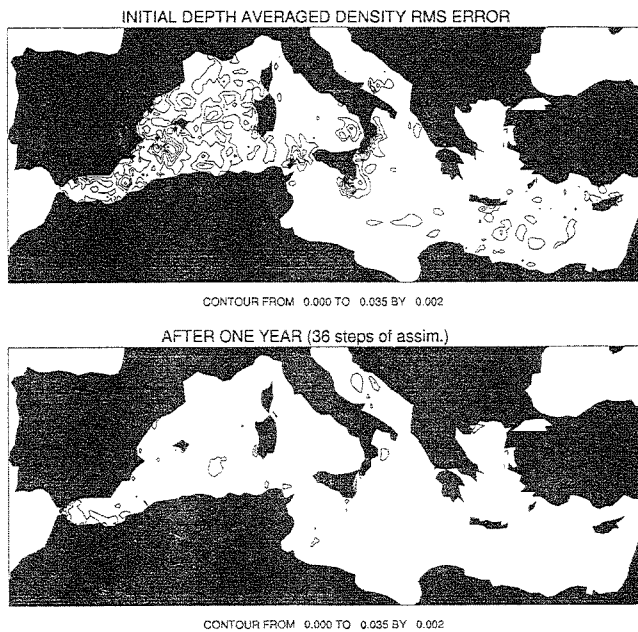


FIG 1: Density error (rms deviation from observations), at the start of the assimilation run and after one year of forward integration which included 36 steps of assimilation.

REFERENCES

- HAINES K., 1994. Dynamics and data assimilation in oceanography, in *Data Assimilation: Tools for Modelling the Ocean in a Global Change Perspective*, edited by P.P. Brasseur and J.C.J. Nihoul, NATO ASI Series, Vol. 1, 19, pp 1-32.
ROUSSENOV V., E. STANEV, V. ARTALE and N. PINARDI, 1994. A seasonal model of the Mediterranean Sea general circulation. In press on the *J. Geophys. Res.*

ESTIMATE OF THE CAPACITY OF NEARSHORE WATERS TO DISPERSE DISCHARGED EFFLUENTS

M. Hany S. ELWANY¹ and John REITZEL²

¹ Center of coastal studies, Scripps Inst. of Oceanography, La Jolla, CA 92093, USA

² Oceanographer, 763A Tourist Park Rd., Halifax, PA 01732, Canada

Full-scale prediction of the spatial-temporal distribution of effluent concentration around a discharge in nearshore waters requires extensive plume-modelling applied to a variety of receiving-water conditions representing the long-term statistics of currents and density stratification. As a preliminary step, it is useful to have a rough estimate of the spatial distribution of vertically-averaged long-term mean concentration of effluent, to predict the ranges along and across the shelf over which the effluent may have appreciable effects on the environment. This kind of rough estimate may be obtained by treating the matter as a diffusion problem, with time-varying diffusivity estimated from the autocovariance functions of long local current records.

In principle, an accurate estimate of long-term diffusivity can only be made from the Lagrangian autocovariance of the velocities of a large set of drifters or dye-marks released from the same point at random intervals over a long period of time (taking the time as zero at the release of each drifter), and not from the Eulerian autocovariance derived from a current-record at a fixed point. In the limit of small times, however, both kinds of autocovariance give the same-time varying diffusivities $K^1 = w_i^2 t$, in which w_i^2 is the long-term variance of velocity in the i th direction. The differences at longer times will not generally spoil a rough estimate.

To deal with a nearshore discharge over a sloping bottom, the diffusion problem may be solved for a space in the shape of a long wedge, bounded by sea-bottom and the surface, with the apex at the shoreline. Taking the discharge as uniform from top to bottom allows the generally unknown vertical diffusivity to drop out of the solution, and gives vertically-averaged concentrations due to the total discharge. The solutions show concentrations directly proportional to discharge rate and inversely proportional to bottom slope, the long-term standard deviation of current velocity, and the square of the distance from shore to the discharge.

An estimate for a continuous cooling water discharge of $1000 \text{ m}^3/\text{sec}$ (ELWANY *et al.*, 1990) into southern California waters, 2.5 km from shore with an average bottom slope of .006, gave long term mean concentrations of discharged water near the shore as about 9 parts per thousand 5 km alongshore from the discharge and about 5 parts per thousand at 10 km alongshore (Fig. 1). The effect of a long-term mean longshore current of 2.9 cm/sec was to displace the whole pattern of concentration about 2 km downcurrent (Fig. 2).

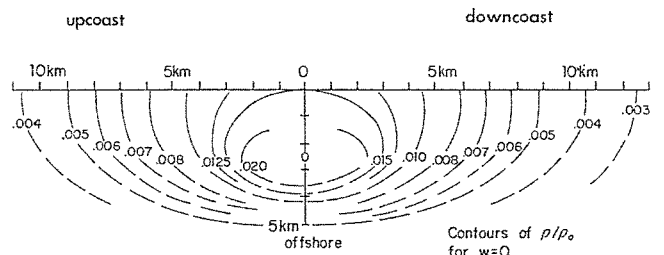


Figure 1. Contours of relative concentration p/p_0 with no mean current ($W = 0$)

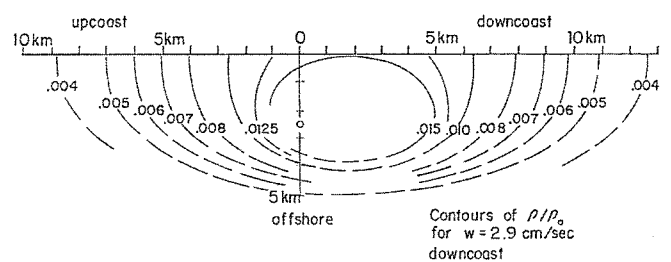


Figure 2. Contours of relative concentrations p/p_0 with for mean current $W = 2.9 \text{ cm/sec}$ downcoast.

REFERENCES

- ELWANY, M. H. S., J. REITZEL and M. R. ERDMAN, 1990. Modification of coastal currents by power plant intake and thermal discharge systems. *Coastal Engineering*, v. 14 p. 359 - 383.