SOME FEATURES OF WATER CIRCULATION AND HYDROLOGICAL STRUCTURE IN THE NORTH-EASTERN PART OF THE LEVANTINE SEA

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² South. Branch of P. P. Shirshov, Inst of Oceano. of the Ac. of Science, Gelendzhik, Russia Hydrophysical observations performed by r/v "Vityaz" in the Levantine Sea were assumed as basin of studies. They were carried out in the region between the island of Cyprus and Syrian coast in winter 1992 and autumn 1993. During the 24th expedition in the course of two hydrophysical surveys (February 15-22 and March 5-8, 1992) 105 stations were made. During the 27th expedition the similar surveys were performed in October 8-13 and 24-28 and comprised 107 stations. The measurements of z, T and S were carried out with Neil Brown probe to the depth of 800-1000 m with the distance between the stations about 10 miles. First of all, it is necessary to note that in severe winter 1992 the whole water column down to the main pycnocline (to the upper boundary of deep waters) was covered by convection. That is why with the increasing of measurement, the well-pronounced multilayer water structure was yet preserved. Therefore in the most upper layer over the seasonal thermocline and under it the water circulation had sometimes large differences including opposite directions of currents. In recent years, a great attention is paid to theoretic and experimental studies of the meandering processes of the Main current in the Mediterranean Sea and eddy making at its lateral boundaries. As is seen from the comparison of dynamic maps estimated relative for the season of dynamic maps estimated relative.

In recent years, a great attention is paid to theoretic and experimental studies of the meandering processes of the Main current in the Mediterranean Sea and eddy making at its lateral boundaries. As is seen from the comparison of dynamic maps estimated relative to 500 dbar depth for several surveys, both in winter (the dynamic map is not given in the summary) and in autumn (Fig. 1) the water circulation in the region under studies is characterised by the strong meandering of the Main Stream, complex eddy structure of currents, their considerable synoptic and seasonal variation. It differs essentially from a classic concept of the main South to North water transport within the polygon. According to the data of four surveys the cyclonic system of circulation with two or several smaller centres is observed to the left of the Main current near the eastern coast of Cyprus. The cyclonic gyre developing to the east of Cyprus is the most stable. The co-ordinates of its centre are $34^{\circ}45^{\circ}$ N and $34^{\circ}30^{\circ}-34^{\circ}45^{\circ}$ E. The forming of large anticyclonic meanders with closed anticyclonic circulation was formed between the eastern coasts of the anticyclonic eddies inside was observed to the right of the Main current. In winter 1992, the anticyclonic circulation was formed between the eastern coasts of the solution 35° E. According to the data of the first winter survey, the anticyclonic dives in part of the polygon and of anticyclonic circulation to 35° E in its southeastern part of the polygon and of anticyclonic circulation is 35° 15 N and 35° 30. E. According to the data of the first winter survey, the general South to North water transport through the most part of the polygon was well traced not only at the surface but also at the intermediate depths. The water circulation at the surface in attumn 1993 (Fig. 1) was characterised by the intensive development of provide date and the surface in attum at the North water transport through the most development of enticyclonic din eatterned general South to North water transport through the most part of the polygon was well traced not only at the surface but also at the intermediate depths. The water circulation at the surface in autumn 1993 (Fig. 1) was characterised by the intensive development of large anticyclonic system in the North and the weak development of anticyclonic circulation near the eastern coast (as compared to winter). That is why the water transport to the North in the north-western part of the polygon turned out to be shifted to the West (to Cyprus), but in the South-East it was traced almost along Syrian coast (with the exception of small local cyclonic eddies near the shore). Dynamic maps for the deeper layers (250500 dbar; Fig. 1b) are essentially different from the water circulation at the surface. The intensive anticyclonic eddy in the northern part of the polygon weakens with depth and is shifting to the North. In the central part of the polygon, the intensive anticyclonic eddy on 34°40° N. This eddy is absent at the surface and is slightly traced only at the depth of 50 m. The maximal development of the eddy was observed between 100 and 250 m where its diameter reached 30 miles. This large anticyclonic, the rising of isolines is observed, but their sharp lowering is observed between 100 and 250 m where its diameter reached 30 miles. This large anticyclonic, the rising of isolines is observed, but their sharp lowering is observed below (to 1000 m). The most interesting are the peculiarities of intermediate water structure of the eddy. They are characterised by the inversions of hydrological water structure of the near shore waters in autumn, are of great interest. These waters formed in the local climatic conditions, may be isolated as a separate surface water structure to the the instructure in the form of a lenge induced and a slainity, weak baroclinic stability. But it may be supposed that these waters in the process of advective transport shore sufface waters with anormaly high salinity (39.50-39.786) and high temper



Figure : Dynamical topography of surfaces 0/500 db(a),250/500 db(b) REFERENCES

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